

EXTENT OF THE NORTHERN BALTIC SEA DURING THE EARLY PALAEOZOIC ERA – NEW EVIDENCE FROM OSTROBOTHNIA, WESTERN FINLAND

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This study reports new evidence of the extent of the northern Baltic Sea during the Cambrian and Ordovician periods. A drillcore (DC304) from the Lappajärvi impact crater, western Finland, and erratics from the surrounding area were studied for acritarchs. The acritarchs from the drillcore were reworked by the explosion but indicate, however, that the Baltic Sea extended beyond the western coast of central Finland during the Lower Cambrian Vergale and the Middle Cambrian Kibartai regional stages as well as during Middle Ordovician transitions from Aseri to Lasnamägi and from Idavere to Jöhvi regional stages.

The provenance of the erratics was the Bothnian Bay, further north than the Lower Palaeozoic deposits previously known in the Bothnian Sea. The age could be determined by their acritarch composition which shows that the sea also covered Central Ostrobothnia during the Lower Ordovician Billingen and early Middle Ordovician Lasnamägi regional stages. The drillcore samples contained also unknown spores. They suggest that there have been sediments and/or a terrestrial flora later than Ordovician in the Lappajärvi area.

Key words: microfossils, acritarchs, spores, sedimentary rocks, impact crater, erratics, Baltic Sea, Paleozoic, Cambrian, Ordovician, Lappajärvi, Alajärvi, Finland

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INTRODUCTION

The previous estimations of the extent of the Baltic Sea during the Early Palaeozoic Era have been based on the drillcore and outcrop material from Denmark (Hansen 1945), Sweden (Hessland 1953–55, Thorslund 1960, Jaanusson 1972, 1982, Lindström 1979, Hagenfeldt 1989a) and Estonia (Õpik 1956, Männil 1966), and they suggest that the sea

has occasionally covered the Bothnian Sea. In Finland only the Söderfjärden drillcore (Fig. 1) has been included in the estimations by Hagenfeldt (1989a).

According to seismic profiles, Early Palaeozoic, i.e. Cambrian and Ordovician, marine deposits cover large areas of the bottom of the Bothnian Sea (the southern part of the Gulf of Bothnia; Winterhalter 1972), and Cambrian deposits the

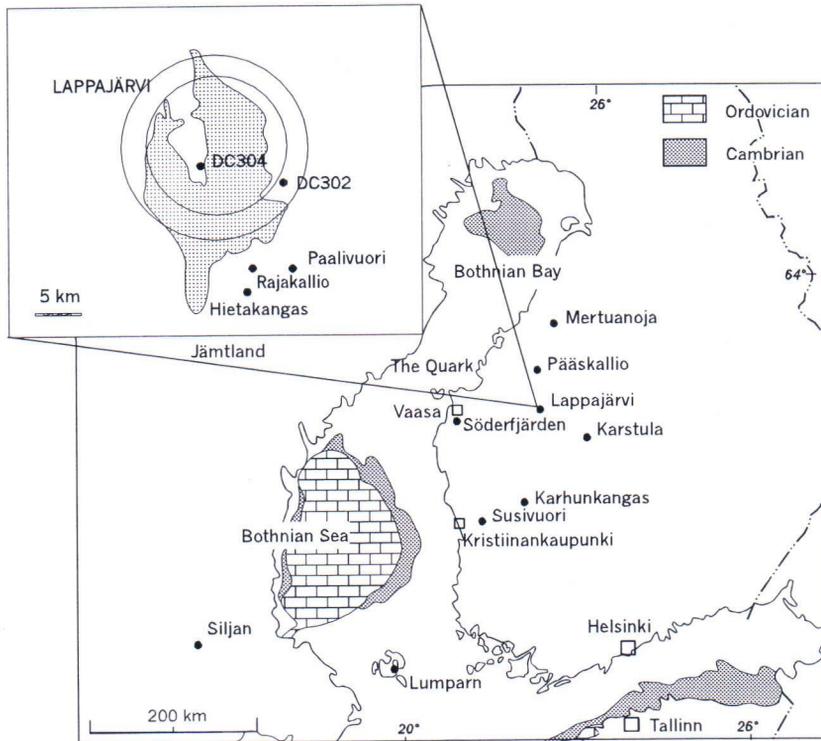


Fig. 1. Location of Lake Lappajärvi and the sites of the erratic boulders.

bottom of the Bothnian Bay (the northern part of the Gulf of Bothnia; Wannäs 1989; Fig. 1). Cambrian fissure fillings occur in the archipelagos of Åland (Bergman 1982) and southwestern Finland (Martinsson 1956, Simonen 1956).

In situ Cambrian sediments cover the Söderfjärden impact crater (Tynni 1978, 1982a, Hagenfeldt 1989b, c), only 100 km west of Lake Lappajärvi (Fig. 1). Sandstone erratics – probably Cambrian in age – have been found in Karstula, Ostrobothnia (Simonen & Kouvo 1955, Tynni 1974) and on the eastern coast of the Bothnian Bay (Uutela 1983). They show the Cambrian sea reached up to 65° N.

Two drillcores from Finngrundén on the Swedish side of the southern Bothnian Sea contain Early Ordovician trilobites (Thorslund & Axberg 1979, Tjernvik & Johansson 1980). Thus far, the only known *in situ* Ordovician sedimentary deposits in Finland are situated at the bottom of

Lumparn Bay, Åland (Kulling 1926, Winterhalter 1972). Uutela (1989) found Cambrian and Ordovician erratic boulders in southwestern Finland up to 62° N. Reworked Ordovician acritarchs have also been found in Quaternary deposits elsewhere in Ostrobothnia (Kujansuu & Uutela 1997, Eriksson et al. in print).

The aim of the study was to investigate Palaeozoic acritarchs from a drillcore at Lappajärvi and from erratic boulders found in Ostrobothnia and to discuss the extent of the northern Baltic Sea during the Early Palaeozoic Era.

THE AREA STUDIED

Lappajärvi is located in Ostrobothnia, Central Finland (Fig. 1). It has a more than 150 year history of geological research: the first hundred years it was interpreted as a volcano (Eskola 1927), sin-

ce 1968 as an impact crater (Svensson 1968). The diameter of the crater is 17 km. The ^{40}Ar - ^{39}Ar age of kárnäite is 77 Ma (Jessberger & Reimold 1980), and the palaeomagnetic age is less than 200 Ma (Pesonen et al. 1984). In drillcore DC302 from the eastern slope of the crater beneath the Quaternary tills, there are 18 metres of sedimentary rocks with Precambrian acritarchs (Uutela 1990).

MATERIAL AND METHODS

The drillcore material (DC304) studied comprises 12 samples of mudstone with acritarchs from the Kärnäsaari island three kilometres southwest of the geometrical centre of the crater. Mudstone at a depth of 32.60–35.20 m is overlain and underlain by suevite and impact breccia material (Fig. 2).

Microfossils of erratic boulders were also studied, as were acritarchs in erratic boulders found in Ostrobothnia. Two boulders of reddish marl and one yellowish Baltic limestone (calcilutite) are from the Hietakangas esker in Höykkylä, Alajärvi (map sheet 2313 08, $x=6992.10$, $y=2485.90$, $z=105$). One greenish marl is from the Rajakallio esker and littoral deposit complex in Höykkylä (2313 08, $x=6996.50$, $y=2488.06$, $z=125$). One reddish marl is from the Paalivuori littoral deposit in Hoisko, Alajärvi (2313 11, $x=6995.10$, $y=2494.60$, $z=130$). All these sites are located 13–17 km southeast of Kärnäsaari. One erratic of reddish marl is from the Susivuori littoral deposit in Kristiinankaupunki (1232 12, $x=6911.00$, $y=1535.00$, $z=120$), 140 km southwest of Lappajärvi, and two from the Pääskallio littoral deposit in Kaustinen (2323 07, $x=7049.30$, $y=2487.45$, $z=81$), 40 km north of Lappajärvi (Fig. 1).

Before the maceration process, samples were treated with 10 % HCl and washed with distilled water to prevent contamination by modern pollen and spores. In the maceration process HCl and HF were used before filtering the residue (Vidal 1988). The sample size of mudstone was ca. 20 g and of limestone 50 g.

The acritarchs studied here are rather poorly preserved. Many specimens of the Ordovician gen-

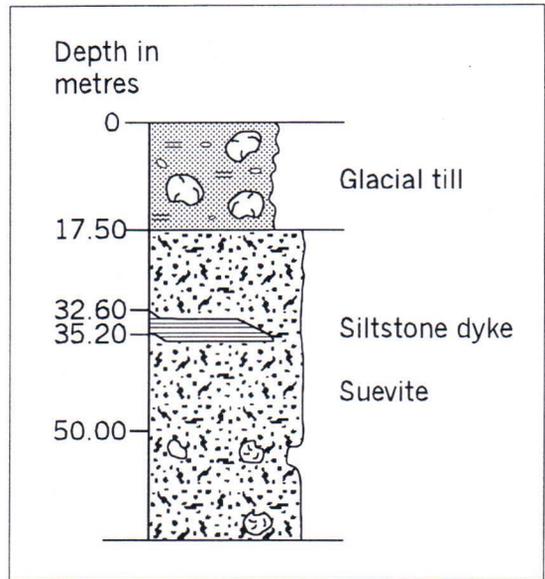


Fig. 2. The sequence of the upper part of the Lappajärvi drillcore DC304.

era *Baltisphaeridium* and *Ordovicidium* are therefore impossible to identify on the species level. Species of the genus *Leiosphaeridia* are excluded from the percentile proportions due to their minimum dating value.

In this study the terms for Estonian regional stages of the Cambrian and Ordovician systems are used because the main reference material of acritarchs is from Estonia. In Estonia, the Ordovician system is subdivided into 19 regional stages, and the Cambrian into eight (Fig. 3).

RESULTS

Acritarchs in the Kärnäsaari drillcore DC304

In the miscellaneous assemblage of microfossils in the Lappajärvi drillcore, 75.4 % are Cambrian and 21.1 % Ordovician acritarchs, in addition to 3.5 % trilete spores of unknown age. The maximum Cambrian specimens occur at a depth of 33.50 m, Ordovician at a depth of 35.00 m and spores at a depth of 33.90 m (Table 1a & b).

ESTONIAN REGIONAL STAGES	BRITISH SERIES			
PORKUNI F _{II}	ASHGILL	UPPER ORDOVICIAN		
PIRCU F _{Ic}				
VORMSI F _{Ib}				
NABALA F _{Ia}	CARADOC	MIDDLE ORDOVICIAN		
RAKVERE E				
OANDU D _{III}				
KEILA D _{II}				
JÕHVI D _I				
IDAVERE C _m				
KUKRUSE C _{II}				
UHAKU C _{Ic}			LLANVIRN	ORDOVICIAN
LASNAMÄGI C _{Ib}				
ASERI C _{Ia}				
KUNDA B _{III}	ARENIG	LOWER ORDOVICIAN		
VOLKHOV B _{II}				
BILLINGEN B _{Ib}	TREMADOC	ORDOVICIAN		
HUNNEBERG B _{Ia}				
VARANGU A _{III}				
PAKERORT A _{II}	UPPER CAMBRIAN			
ULGASE	MIDDLE CAMBRIAN			
PANERIAI	LOWER CAMBRIAN			
KYBARTAI				
RAUSVE				
VERGALE				
TALSY				
LONTOVA				
ROVNO				

Fig. 3. Cambrian and Ordovician Estonian regional stages. References: Mens et al. (1987) and Nõlvak (1997).

Kärnäsaari mudstone includes 1.32–5.35 % CaCO₃ (Table 1a). A large amount of CaCO₃ does not indicate a large number of Ordovician acritarch species (correlation coefficient R=0.448).

Cambrian species

Living ranges of Cambrian species are presented in Table 2; 8.9 % of species are restricted to the Early Cambrian and 61.0% of species have a living range from the Early to Middle Cambrian.

Filisphaeridium molliculum (0.6 %; Fig. 4a) disappears in the Early Cambrian Vergale Regional Stage, where a living range of *Helosphaeridium coniferum* (0.6 %; Fig. 4b) is restricted. In addition to the southern Baltic Sea, *F. molliculum* is recorded at Söderfjärden (Hagenfeldt 1989b), *H. coniferum* is reported in Poland, Scotland, Spitzbergen and Canada (Moczydlowska 1991 and references therein), but this is the first time it has

been recorded in the northern Baltic Sea. *Liepainna plana* and *Micrhystridium notatum* (totally 0.1 %) appear in the upper part of the Rausve Regional Stage during the Early Cambrian. These species are also reported at Söderfjärden (Hagenfeldt 1989b).

The only species restricted to the Middle Cambrian Kibartai Regional Stage is *Baltisphaeridium pseudofoveolatum* (0.06 %; Fig. 4c). It has been earlier reported in Sweden, Latvia, Lithuania and the Kaliningrad Region (Hagenfeldt 1989b and references therein).

Ordovician species

Living ranges of Ordovician species are presented in Table 3; 29 % of species have a long living range from the Early to Late Ordovician, 9 % of species have lived from the Early Ordovician to the early Middle Ordovician and 12.2 % of species from the late Middle Ordovician to the Late Ordovician (Uutela & Tynni 1991 and references therein).

Based on the preliminary results from the Mänamaa material, Hiiumaa, the living range of *Peteinosphaeridium eximum* (Fig. 4d) is restricted in the Volkhov (B_{II}) to Aseri (C_{Ia}) regional stages in Estonia. In the early Middle Ordovician Lasnamägi Regional Stage (C_{Ib}) *Baltisphaeridium dasos*, *Ordoviciidium grotlingboensis* and *O. nudum* appear. Later in the Middle Ordovician, i.e. at the beginning of the Jõhvi Regional Stage (D_I), appears *Cheleutochroa venosa* (Fig. 4e). Other species with living ranges restricted to the late Middle Ordovician are *Baltisphaeridium trabeculaespiniae*, *Ordoviciidium elegantulum*, *Verychium irroratum*, *V. reductum* and *V. trisulcum*.

Unknown spores

The material includes two previously unknown types of spores.

Spore 1. The triangular amb has straight or slightly convex sides. Triletic sutures are straight, usually extending to the periphery of the spore. The thin exine (0.5 µm) is scabrate or microgranulate. The length of side varies 20–30 µm, mean

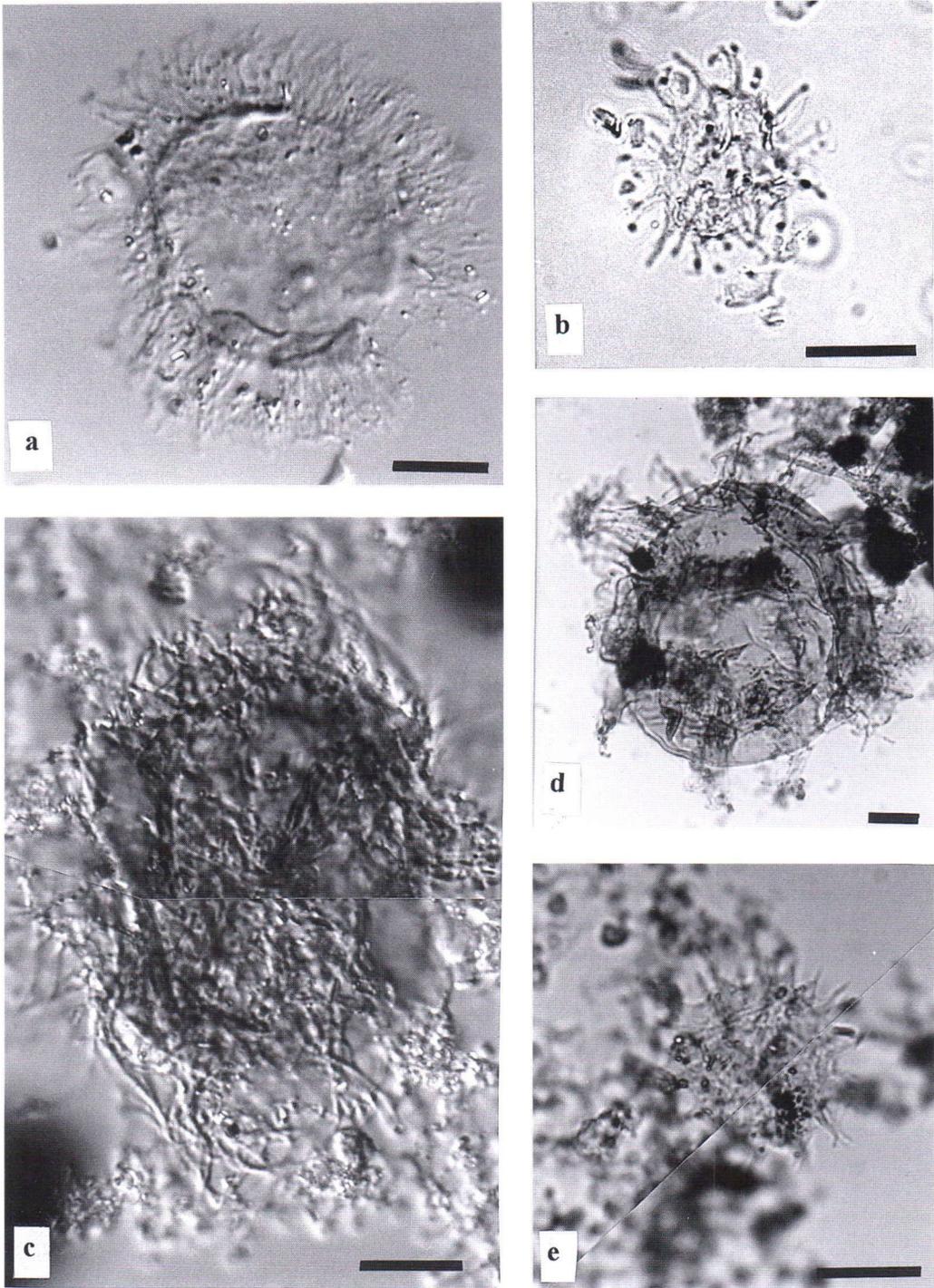


Fig. 4. Cambrian and Ordovician acritarchs in the Lappajärvi DC304: *Filisphaeridium molliculum* (a), *Helosphaeridium coniferum* (b), *Baltisphaeridium pseudofoveolatum* (c), *Petinosphaeridium eximum* (d) and *Cheleutochroa venosa* (e). Bar equals 10 μ m.

is 25 μm . A total of 74 specimens were recorded. Spore 1 resembles the modern spore of *Pteridium* (Figs. 5a-d).

Spore 2. The amb is triangular, sides are straight. Triletic sutures are straight, usually extending to the periphery of the spore. The rather thick exine (1.5 μm) is microgranulate. The length of side is 22 μm . Only one specimen is recorded at the depth of 34.50 m. Spore 2 resembles the modern spore of *Sphagnum* (Figs. 5e-f).

The oldest spores from the Precambrian deposits in Russia and Latvia have been reported by Volkova (1976), but instead of a defined trilete suture they have trilete folds. The terrestrial Ordovician spores reported by Vavrdová (1984) and the Silurian miospores reported by Richardson (1996) differ greatly from the spores 1 and 2. Among the present assemblage 31 % of acritarchs have a living range up to the Lower Silurian or later. Because real Silurian indicator species are missing, it seems improbable that Silurian deposits also covered the Lappajärvi area. Richardson (pers. comm.) thinks the spores are possibly post-Frasnian (Upper Devonian) or later.

No pollen grains were found. This speaks against the contamination by modern material since pollen grains are more abundant than spores in modern deposits. Therefore it is probable that the spores were deposited when sporophytes had already developed but before the appearance of pollen producing plants, i.e. during the Carboniferous.

The Lappajärvi area before the Quaternary ice age

According to the results, the area where Lappajärvi is situated was covered with Lower Palaeozoic deposits before the meteorite impact. The lowermost deposits consisted of Lower Cambrian Vergale deposits overlying Middle Cambrian Kibartai deposits. Middle Cambrian Rausve deposits are possible, but not verified here. These were mixed during the explosion and fell down into a fissure filling of suevite.

The Ordovician acritarchs suggest that the sea covered the Lappajärvi area at least during the

transitions from Aseri to Lasnamägi and from Idaverre to Jöhvi regional stages, Middle Ordovician.

It seems that Precambrian deposits covered only the eastern side of Lappajärvi because the present assemblages lack the Precambrian specimens found by Uutela (1990) on the rim area of the Lappajärvi impact crater.

Although the two types of spores remain unidentified, they indicate that during an unknown period after the Ordovician, sediments containing spores have been deposited in the Lappajärvi area.

Acritarchs of the Hietakangas erratic boulder

Among the erratic boulders only the yellowish grey Baltic limestone (calclutite) from the Hietakangas esker yielded 44 rather poorly preserved acritarch specimens. The assemblage is typical of the Baltic area; most of species belong to genera *Baltisphaeridium* and *Ordovicidium* (Table 4). *Baltisphaeridium esthonicum* (Fig. 6a) disappears in the Lasnamägi Regional Stage (C₁b), when *B. parvigranosum* (Fig. 6b), *Ordovicidium grotlingboensis* (Fig. 6c) and *O. nudum* (Fig. 6d) appear. Thus the age of the Hietakangas boulder is the Lasnamägi Regional Stage, early Middle Ordovician.

Other limestone erratics

Other limestone erratics contained no acritarchs. Oxidation has destroyed organic-walled fossils in reddish marl. Conodonts are better for dating this type of limestone (Löfgren 1985), but a larger sample size, more than 500 g, would be needed. In the northern Baltic Sea the Finngrunden reddish marl is Arenig in age (Tjernvik & Johansson 1980). It correlates with the Estonian regional stages of Billingen, Volkhov and the beginning of Kunda in Estonia. Based on lithological features the Hietakangas, Paalivuori, Pääskallio and Susivuori boulders of reddish marl are of the same age.

The age determination of the Rajakallio greenish marl boulder is complicated and remains unsolved. Deposits of greenish marl are found in the Rapla drillcore in the Middle Ordovician Jöhvi and Oandu regional stages (Põlma 1972, Uutela &

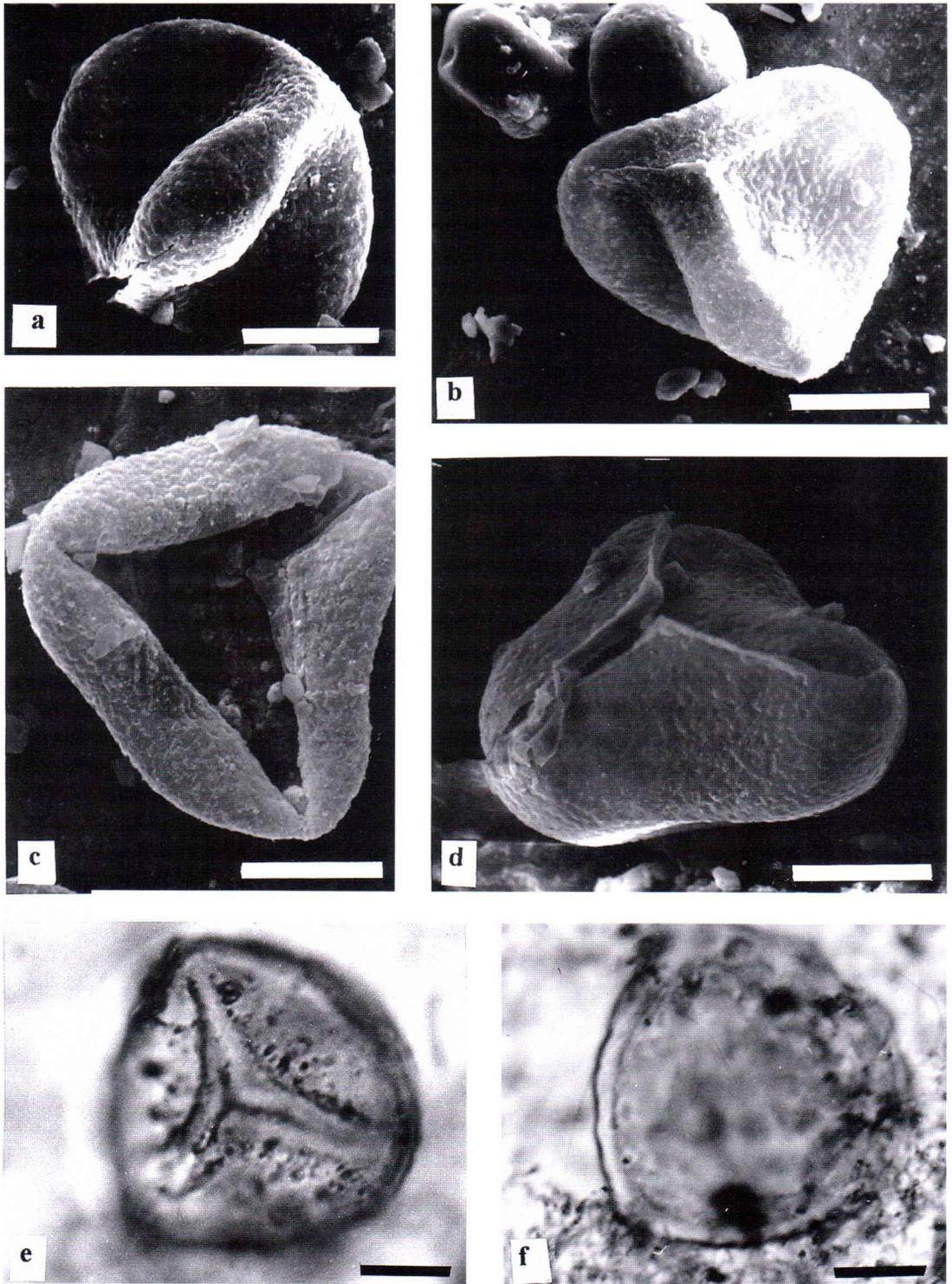


Fig. 5. Unknown spores in the Lappajärvi DC304. Spore 1 (a-d) and spore 2 (e-f). Bar equals 10 µm.

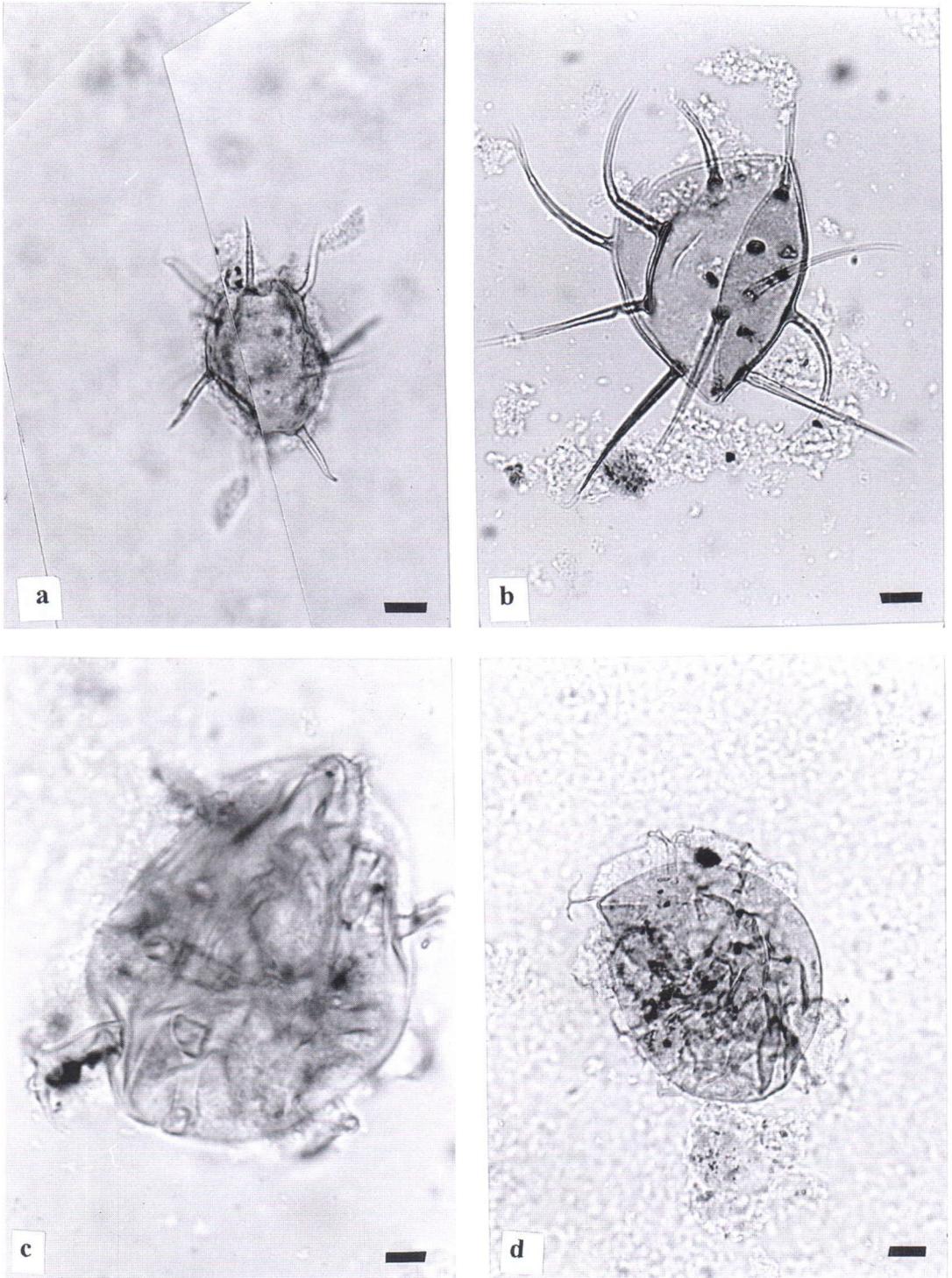


Fig. 6. Ordovician acritarchs in the Hietakangas erratic boulder. *Baltisphaeridium esthonicum* (a), *B. parvigranosum* (b), *Ordoviciidium grotlingboensis* (c) and *O. nudum* (d). Bar equals 10 μm .

Tynni 1991) and in the Männamaa drillcore from the Jõhvi to Vormsi regional stages, Middle to Upper Ordovician (Uutela in prep.). Greenish marl was also deposited in Estonia during the Early Silurian Juuru Regional Stage (Põlma 1972, Uutela & Tynni 1991, Uutela in prep.). According to estimations by Männil (1966), greenish marl was deposited during the Lasnamägi to Idavere regional stages in the northern Baltic Sea.

Provenance of the Ordovician limestone erratics

As mentioned above, there are Ordovician *in situ* deposits at the bottom of the Bothnian Sea (Fig. 1). The northernmost border reaches 62°40' N, like at Kristiinankaupunki, and the westernmost side reaches near 20°00' E (Winterhalter 1972). Because of the glacial movement from the northwest, this area is not a possible provenance of the Ostrobothnian boulders.

The meteorite impacted a large area of sedimentary deposits in the surroundings of the present lake of Lappajärvi. Because of the lack of acritarchs, it is impossible to verify if the provenance of the boulders of reddish marl from Hietakangas and Paalivuori, as well as the Rajakallio greenish marl boulder are near this area. The location of the Pääskallio boulder indicates that the provenance of boulders of reddish marl found there is more northern than the Lappajärvi area (Fig. 1). The Hietakangas boulder of calcilutite seems to be from the similar deposits that covered the Lappajärvi area. In the Lappajärvi – Kaustinen area the latest advancement of the continental ice has taken place from the northwest (Mölder 1948, Geological Survey of Finland 1986).

The provenance of the reddish marl from Susivuori is possibly different from that of the other boulders, because in the Kristiinankaupunki area the continental ice flow came from the north (Mölder 1948, Geological Survey of Finland 1986).

Reworked Ordovician acritarchs from Kauhajoki may be from the same provenance, i.e. from the Quark or the Bothnian Bay (Kujansuu & Uutela 1997). Reworked Ordovician acritarchs found

in Eemian deposits at Mertuanoja (Eriksson et al. in print) suggest that there may have been Lower Palaeozoic deposits as far north as at the bottom of the Bothnian Bay. The acritarch assemblage of Mertuanoja indicates deposits of the Lasnamägi Regional Stage (*Baltisphaeridium parvigranulosum*).

DISCUSSION AND CONCLUSIONS

The extent of the Cambrian sea in the northern Baltic Sea

For Hessland (1955), the extent of the northern part of the Baltic Sea remained an unsolved question. According to Öpik (1956), the sea covered the Bothnian Sea and the Bothnian Bay at the beginning of early Cambrian. This occurred during the Vergale Regional Stage and at the beginning of the Rausve Regional Stage (late Early Cambrian), but not as late as at the end of the Rausve Regional Stage (Hagenfeldt 1989a). The results of the acritarch analysis in this study support these results with a wider eastern extent, however (Fig. 7).

Hansen (1945) and Thorslund (1960) suggested that the early and late Middle Cambrian Baltic Sea reached only to southeastern Sweden, Denmark and western Norway. The presence of *Baltisphaeridium pseudofoveolatum* in this study – although only one specimen – indicates that the sea did reach up to the Lappajärvi area during the early Middle Cambrian Kibartai Regional Stage, a similar result as reported by Hagenfeldt (1989a), except a wider eastern extent (Fig. 7). Late Middle Cambrian acritarchs (the Paneriai Regional Stage) were not found in this material. This supports the theory of Hagenfeldt (1989a) that the sea covered the Bothnian Sea, but not the western coast of Finland. Thus, the Lappajärvi area was dry land during the late Middle Cambrian.

The extent of the Ordovician sea

At least two different types of Ordovician acritarch assemblages can be distinguished in the Lappajärvi

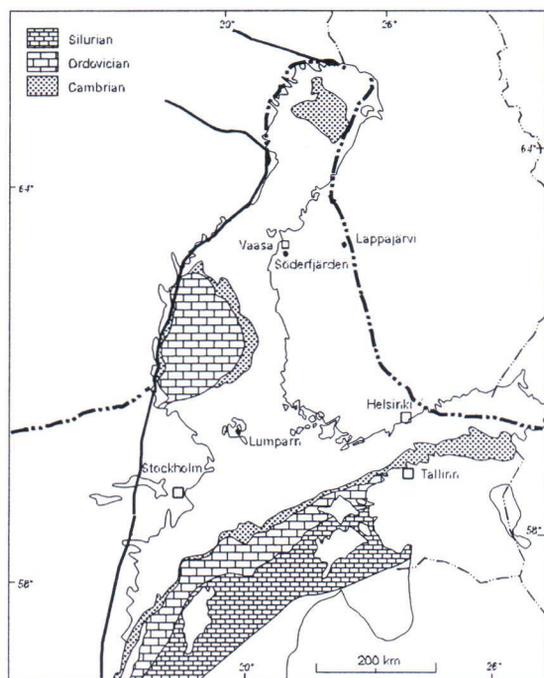


Fig. 7. Extent of the Lower (whole line) and early Middle Cambrian Baltic Sea (dashed line), the Swedish side of the sea is according to Hagenfeldt (1989a).

vi material. The older one is from the transition of the Aseri (C_{1a}) to Lasnamägi (C_{1b}) regional stages (early Middle Ordovician). The more recent one is from the transition of the Idavere (C_{III}) to Jõhvi (D_1) regional stages (middle of Middle Ordovician). Age determinations of erratic boulders indicate that during the Billingen – Kunda (reddish marl) and Lasnamägi (yellowish calcilutite) regional stages, the Ordovician sea covered the northern Bothnian Sea.

According to Thorslund (1960), the Early Ordovician (Early Tremadocian, i.e. Pakerort Regional Stage) sea extended only to the southern coast of Finland. Männil (1966) estimates that during the Pakerort Regional Stage (A_{II}) quartz-rich sand was deposited in the central Baltic Sea. Deposits of *Dictyonema* shale *in situ*, Pakerort in age in the Siljan district, Central Sweden (Jaanusson 1982) speaks against both theories.

According to Männil (1966), during the Late Tremadoc (i.e. Varangu Regional Stage) and Arenig to early Llanvirn (i.e. Hunneberg, Volhov, Kunda and Aseri regional stages; B_{II} – C_{1a}) only the eastern Bothnian Sea was covered with oceanic waters and yellowish grey Baltic limestone (calcilutite), and reddish and greenish marls were deposited. *In situ* deposits of these ages in the Siljan district and further north in Jämtland (Jaanusson 1982) also contradicted Männil's theories. According to Lindström (1979), during the Arenig the sea covered wide areas of southern Finland, and Thorslund (1960) interpreted the sea as extending north to western Finnish Lapland. The present material, especially the erratic boulders of reddish marl, suggest that the Early Ordovician sea reached out to the Bothnian Bay (Fig. 8) and thus supports the theories of Thorslund (1960) and Lindström (1979). One of the Ordovician transgressions in Europe took place during the transition of Tremadoc to Arenig (Fortey 1984), in the Estonian regional stages it took place during the Varangu and Billingen regional stages. Based on this evidence, the reddish marl is deposited during the Billingen Regional Stage, instead of Billingen – Kunda (cf. p. 56).

During the Middle Ordovician Lasnamägi, Uhaku, Kukruse and Idavere regional stages (C_{1b} – C_{III}) greenish marl was deposited in the Bothnian Sea before the regression in the eastern Baltic Sea during the Jõhvi, Keila and Oandu regional stages (D_1 – D_{III}) (Männil 1966). According to Jaanusson (1982), the sea covered Dalarna all the period from Lasnamägi to as late as Porkuni, from middle to late Ordovician. The second Ordovician transgression in Europe took place during Llanvirn (Fortey 1984), i.e. during the transition from Aseri to Lasnamägi. The results of the present study (Fig. 9) support this view.

During the Caradoc, one of the widest Ordovician transgressions of the world also covered most of the epeiric seas (Fortey 1984). It was in the beginning of this transgression (Idavere and Jõhvi regional stages) that the Lappajärvi area was covered by the sea. This view is in accordance with the results of this study.

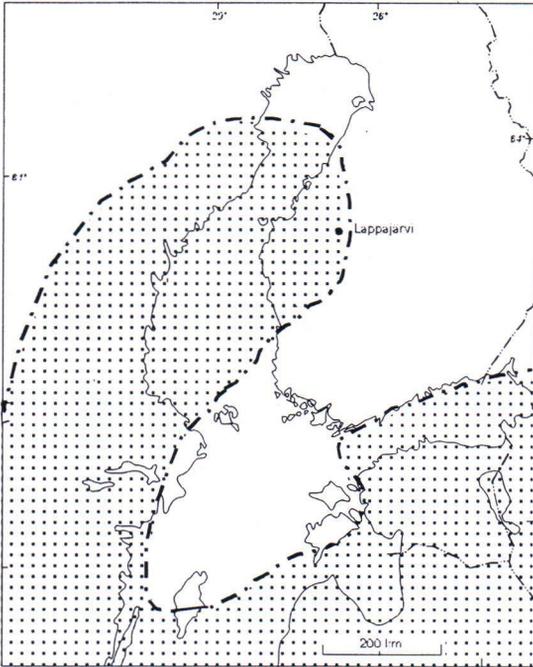


Fig. 8. Extent of the Lower Ordovician Baltic Sea (Billingen Regional Stage).

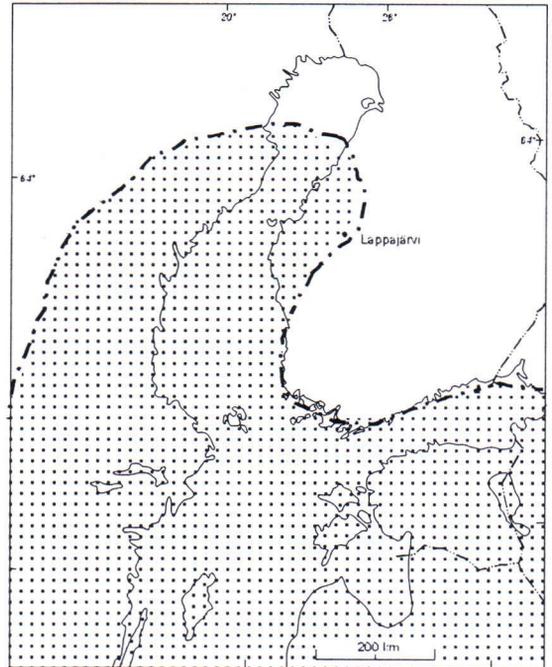


Fig. 9. Extent of the Middle Ordovician Baltic Sea (Aseri/Lasnamägi transition and Idavere/Jõhvhi transition).

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ript and Phil.lic. Tuija Jantunen drew some of the figures. Professors Joakim Donner and Pentti Alhonen refereed the manuscript. The English language was checked by Margot Whiting. I thank all of them cordially.

Table 1a. The total number of Cambrian acritarch specimens, and percentual amounts of CaCO₃ in mudstone in the Kärnäsaari drillcore DC304, Lappajärvi.

Depth (m)	32.60	32.80	33.15	33.30	33.50	33.70	33.90	34.10	34.50	34.70	35.00	35.20	Σ
Acritarch species													
<i>Archaeodiscina umbonulata</i>		2		4	3		2	3	3	1	7	3	28
<i>Baltisphaeridium pseudofoveolatum</i>								1					1
<i>Cymatiosphaera postii</i>				1	12	6	1	3	1		4	8	36
<i>Filisphaeridium cerinum</i>				3	9		1		2		2	1	18
<i>F. dubium</i>				1	2		5		2		2	7	19
<i>F. molliculum</i>				2	4		1		2		1		10
<i>F. lanatum</i>					2			3	1			2	8
<i>F. tornatum</i>		2		6	40	2	15	6	19	11	38	33	170
<i>Granomarginata prima</i>				17			12					7	36
<i>G. squamacea</i>				4	8	1	4	5	6	3	2	2	35
<i>Helosphaeridium coniferum</i>				1	7			1					9
<i>H. lubomlense</i>					1								1
<i>Leiomarginata simplex</i>				23	83	32	13	71	73	18	37	70	420
<i>Leiosphaeridia gregalis</i>				13	11	8	16	3	10	1	11	8	81
<i>Liepaina plana</i>				1	1							1	3
<i>Lophosphaeridium tentativum</i>		5			16	3	4	22	22	3	35	37	147
<i>Micrhystridium dissimilare</i>					2	1			1				4
<i>M. notatum</i>				1	2								3
<i>M. obscurum</i>				1									1
<i>Pterospermella velata</i>					1		2						3
<i>Skiagia ciliosa</i>			1	31	65	10	42	29	13	3	37	54	285
<i>S. compressa</i>				20	56	4	11	16	1	2	15	25	150
<i>S. insigne</i>								1					1
<i>S. orbiculare</i>				5	29		2	2	1		2	4	45
<i>S. ornata</i>				8	19		3				2		32
<i>Tasmanites bobrowskiae</i>		5			2		8	2			1	4	22
<i>T. tenellus</i>		1				1	8	2					12
<i>T. volkovae</i>		7	2	8	12	3	2		2		5	11	52
<i>T. sp.</i>		3			1						1		5
Cambrian species	0	25	3	150	388	83	140	170	159	42	202	277	1639
CaCO₃ (%)	2.15	4.98	2.19	3.77	1.32	2.02	3.33	3.97	3.18	2.65	5.35	2.80	35.7
<i>Leiosphaeridia</i> spp.	7	63	12	173	306	133	238	205	368	87	486	164	2242

Table 1b. The total number of Ordovician acritarch specimens, and unknown spores in the Kärnäsaari drill-core DC304, Lappajärvi.

Acritarch species	Depth (m)	32.60	32.80	33.15	33.30	33.50	33.70	33.90	34.10	34.50	34.70	35.00	35.20	Σ
<i>Baltisphaeridium accinctum</i>						1								1
<i>B. anneliae</i>					1	1	2		2	1				7
<i>B. constrictum</i>						1	1	1						3
<i>B. dasos</i>						1								1
<i>B. flexuosum</i>					2									2
<i>B. hirsutoides</i>						6	2	4	1		1	4	7	25
<i>B. ingerae</i>									1					1
<i>B. latiradiatum</i>													1	1
<i>B. longispinosum</i>						2				1		1	2	6
<i>B. microspinosum</i>							1							1
<i>B. multitechinatum</i>									1					1
<i>B. multipilosum</i>						1								1
<i>B. nanninum</i>			2			1		2	2					7
<i>B. oligopsakium</i>												1		1
<i>B. parvigranosum</i>									1					1
<i>B. pauciverrucosum</i>								1						1
<i>B. pseudocalicispinum</i>								1	1					2
<i>B. pustulatum</i>						1								1
<i>B. trabeculaespiniae</i>					1									1
<i>B. tranvikensis</i>			4								3	1		8
<i>B. sp./Ordoviciidium sp.</i>	3	9	4	2	2			2	3	1		3	6	35
<i>Cheleutochroa venosa</i>													1	1
<i>Comasphaeridium filosum</i>			1			1		3	1					6
<i>Filisphaeridium henryi</i>					1									1
<i>Goniosphaeridium mochtensis</i>			2		3	1	2	2		1	2		1	14
<i>G. sp.</i>					5		1							6
<i>Lophosphaeridium citrinipellatum</i>								1		4	1	1		7
<i>L. disparipelliculum</i>					2									2
<i>L. papillatum</i>			3		6	5		4	12					30
<i>L. pilosum</i>												1		1
<i>Martinisphaeridium bifurcatum</i>								1	1					2
<i>M. bipalmatum</i>										1		1		2
<i>M. parvipinnatum</i>						1								1
<i>M. raplaense</i>				1								1	1	3
<i>Micrhystridium nannacanthum</i>										6		3		9
<i>M. stellatum</i>						1		1				1		3
<i>Multiplicisphaeridium alloteai</i>								2	7	5			2	16
<i>M. fissile</i>					2	7			1	3			3	16
<i>M. irregulare</i>						3								3
<i>Ordoviciidium elegantulum</i>			5			3	1		2			1		12
<i>O. grotlingboensis</i>						1	1					1	1	4
<i>O. heteromorphicum</i>			2		1	1		1				2		7
<i>O. nudum</i>								3				2		5
<i>Orthosphaeridium densiverrucosum</i>					1					1				2
<i>O. insculptum</i>						2	1	3						6
<i>O. vibrissiferum</i>													1	1
<i>O. sp.</i>												1		1
<i>Peteinosphaeridium eximum</i>									2					2
<i>Polygonium pellicidum</i>						4	1		1	2	1	1	4	14
<i>Rhopaliophora sp.</i>						1								1
<i>Tasmanites verrucosus</i>					1			2					2	5
<i>Tylotopalla varipinnosum</i>								1				1		2
<i>Veryhachium irroratum</i>					1	4		15	5	1		17	2	45
<i>V. lairdi</i>						1								1
<i>V. reductum</i>					1				2		3	2		8
<i>V. trispinosum</i>					3	5	4	5	18	12	3	5		55
<i>V. trisulcum</i>					10			5		9	4	15	13	56
<i>V. sp.</i>			1											1
Ordovician species		3	29	4	44	57	16	61	64	48	18	66	47	457
Spore sp. 1					7	13	5	18	11	6	3	4	7	74
Spore sp. 2										1				1

Table 2. Living ranges of the Lower and Middle Cambrian acritarchs in the Baltic area. (References: Hagenfeldt 1989b, c, Mozydlowska 1989 with their references). Legend: ⚡ = a living range to earlier than the Talsy Regional Stage. Upper Cambrian assemblage in the Baltic area differs greatly from the earlier ones (cf. Paalits 1995).

Age	Lower Cambrian			Middle Cambrian	
	Talsy	Vergale	Rausve	Kibartai	Paneriai
Acritarch species					
<i>Archaeodiscina umbonulata</i>	⚡	*	*	*	
<i>Baltisphaeridium pseudofoveolatum</i>				*	
<i>Cymatiosphaera postii</i>	*	*	*	*	*
<i>Filisphaeridium cerinum</i>	⚡	*	*		
<i>F. dubium</i>	*	*	*		
<i>F. molliculum</i>	*	*			
<i>F. lanatum</i>	⚡	*	*	*	
<i>F. tornatum</i>	⚡	*	*	*	*
<i>Granomarginata prima</i>	⚡	*	*	*	*
<i>G. squamacea</i>	⚡	*	*	*	*
<i>Helosphaeridium coniferum</i>	*	*			
<i>H. lubondense</i>	*	*	*		
<i>Leiomarginata simplex</i>	⚡	*	*	*	*
<i>Leiosphaeridia gregalis</i>	*	*	*	*	*
<i>Liepaina plana</i>			*	*	*
<i>Lophosphaeridium tentativum</i>	⚡	*	*	*	*
<i>Michystridium dissimilare</i>	*	*	*	*	*
<i>M. notatum</i>			*	*	*
<i>M. obscurum</i>	*	*	*	*	*
<i>Pterospermella velata</i>	⚡	*	*	*	*
<i>Skiagia ciliosa</i>	*	*	*	*	
<i>S. compressa</i>	*	*	*	*	
<i>S. insigne</i>	*	*	*	*	*
<i>S. orbiculare</i>	*	*	*		
<i>S. ornata</i>	*	*	*		
<i>Tasmanites bobrowskae</i>	⚡	*	*	*	*
<i>T. tenellus</i>	⚡	*	*		
<i>T. volkovae</i>	*	*	*	*	*

Table 3. Living ranges of the Ordovician acritarchs in the Baltic area. Legend: ☞ = a living range extending to the Vormsi Regional Stage (F,b) or later. References: Uutela and Tynni (1991) and references herein; ++ preliminary results from the Männamaa material, Hiiumaa, Estonia.

Estonian regional stages	B _{II}	B _{III}	C _{1a}	C _{1b}	C _{1c}	C _{II}	C _{III}	D _I	D _{II}	D _{III}	E	F _{1a}
Acritarch species												
<i>Baltisphaeridium accinctum</i>	*	*	*	*	*	*						
<i>B. anneliae</i>		*	*	*	*	*						
<i>B. constrictum</i> ++				*	*	*	*					
<i>B. dasos</i>				*	*	*	*	*	*	*	*	☞
<i>B. flexuosum</i>		*	*	*	*	*						
<i>B. hirsutoides</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>B. ingerae</i>	*	*	*	*	*	*						
<i>B. latiradiatum</i>		*	*	*	*	*	*	*	*			
<i>B. longispinosum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>B. microspinosum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>B. multiechinatum</i>	*	*	*	*	*	*	*					
<i>B. multipilosum</i>	*	*	*	*	*	*						
<i>B. nanninum</i>		*	*	*	*	*	*	*	*	*		
<i>B. oligopsakium</i>		*	*	*	*	*	*	*	*	*	*	☞
<i>B. parvigranosum</i>				*	*	*	*	*	*	*	*	☞
<i>B. pauciverrucosum</i>	*	*	*	*	*	*	*	*	*	*	*	
<i>B. pseudocalicispinum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>B. pustulatum</i>			*	*	*	*	*	*	*	*	*	☞
<i>B. trabeculaespiniae</i>							*	*	*	*	*	☞
<i>B. tranvikensis</i>	*	*	*	*	*							
<i>Cheleutochroa venosa</i>								*	*	*	*	☞
<i>Comasphaeridium filosum</i>	*	*	*	*	*	*	*	*	*			
<i>Filisphaeridium henryi</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>Gontosphaeridium mochtiensis</i>		*	*	*								
<i>Lophosphaeridium citrinipeltatum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>L. disparipelliculum</i>				*	*	*	*	*	*	*	*	☞
<i>L. papillatum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>L. pilosum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>Martinsphaeridium bifurcatum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>M. bipalmatum</i>				*	*	*	*	*	*	*	*	☞
<i>M. parvipinnatum</i>		*	*	*	*	*	*	*	*	*	*	☞
<i>M. raplaense</i>		*	*	*	*	*	*	*	*	*	*	☞
<i>Michrystridium nannacanthum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>M. stellatum</i>		*	*	*	*	*	*	*	*	*	*	☞
<i>M. varipinnosum</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>Multiplicisphaeridium alloteaui</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>M. fissile</i>	*	*	*	*	*	*	*	*	*	*	*	☞
<i>M. irregulare</i>		*	*	*	*	*	*	*	*	*	*	☞
<i>Ordoviciidium elegantulum</i>					*	*	*	*	*	*	*	☞
<i>O. grotlingboensis</i>				*	*	*	*	*	*	*	*	☞
<i>O. heteromorphicum</i>			*	*	*	*	*	*	*	*	*	☞
<i>O. nudum</i>				*	*	*	*	*	*	*	*	☞
<i>Orthosphaeridium densiverrucosum</i>			*	*	*	*	*	*	*	*		
<i>O. insculptum</i>			*	*	*	*	*	*	*	*	*	☞
<i>O. vibrissiferum</i>			*	*	*	*	*	*	*	*	*	☞
<i>Peteinosphaeridium eximum</i> ++	*	*	*									
<i>Polygonium pellicidum</i>		*	*	*	*	*	*	*	*	*	*	*
<i>Tasmanites verrucosus</i>					*	*	*	*	*	*	*	☞
<i>Veryhachium irroratum</i>					*	*	*	*	*	*	*	☞
<i>V. lairdi</i>		*	*	*	*	*	*	*	*	*	*	☞
<i>V. reductum</i>					*	*	*	*	*			
<i>V. trispinosum</i>		*	*	*	*	*	*	*	*	*	*	
<i>V. trisulcum</i>					*	*	*	*	*			

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