RICHTERITE AND ACTINOLITE FROM THE SIILINJÄRVI CARBONATITE COMPLEX, FINLAND

KAUKO PUUSTINEN

PUUSTINEN, KAUKO 1972: Richterite and actinolite from the Siilinjärvi carbonatite complex, Finland. Bull. Geol. Soc. Finland 44, 83-86.

Chemical, optical and X-ray data are given for the richterite in the glimmerite and actinolite in the synite of the Siilinjärvi carbonatite complex.

Kauko Puustinen, Technical University, Dept. of Mining and Metallurgy, 02150 Otaniemi, Finland.

Introduction

The Siilinjärvi carbonatite complex is situated some 20 km to the north of the city of Kuopio, Eastern Finland, at lat. 63°08' N and long. 27°44' E. It forms a roughly tabular, subvertical body, some 16 km long and up to 1.5 km wide. A detailed description of the complex has been given by Puustinen (1971).

The sequence of emplacement began with the intrusion of the ultramafic phase (glimmerite), followed by the intrusion of syenite. The carbonatite proper (sövite) was intruded into the glimmeritic rocks.

Glimmerite is a quartz- and feldspar-free rock which consists of phlogopite, alkali amphibole, apatite and some calcite. It is medium grained with color varying from red-brown to black. The phlogopite is sometimes reddish-brown and shows an anomalous reverse pleochroism (tetraferriphlogopite). Zircon is often found as an accessory mineral in connection with the amphibole-rich types of glimmerite.

Syenite consists mainly of microcline, amphibole and pyroxene. Albite, quartz and biotite are found as accessories.

Carbonatite proper forms mixed rocks together with the glimmerite, the amount of amphibole in the carbonatite being, however, neglible.

Occurrence

Amphibole is very common in the glimmerite. Locally its amount may be up to 50 % by volume but usually is less than 15 %. It represents a primary crystallization and does not show any sign of alterations.

In hand specimens the amphibole occurs in many cases as euhedral crystals which measure up to 20 cm. The color is usually dark green. Mineralogical data for the alkali amphibole of glimmerite shows that it is related to richterite.

The amphibole from syenite is actinolite. The total amount of amphibole and pyroxene in syenite is usually less than 25 % by volume. The pyroxene is an aegirine-augite.

To the naked eye the differences between the Siilinjärvi actinolite and richterite are very small and therefore the nature of an amphibole is determined only by detailed studies.

X-ray crystallography

The unit cell dimensions (Table 1) were determined from zero-level Buerger precission photographs using MoK α -radiation (Nos 1 and 3) and from Weissenberg photographs using CuK α -radiation (No 2). The β -angle of No 2 was measured from a precession photograph. The films were calibrated with the reflections of an oriented silicon crystal. The accuracy obtained corresponds to probable errors of ± 0.010 , ± 0.020 , ± 0.005 Å and 0°05' for a₀, b₀, c₀ and β respectively.

TABLE 1

Physical properties and unit cell dimensions of richterite from glimmerite and actinolite from syenite, Siilinjärvi carbonatite complex

	1.	2.	3.
			1
a	1.624	1.618	1.622
β	1.642	1.630	1.638
	1.650	1.640	1.646
<i>y</i> — <i>a</i>	0.024	0.022	0.024
2Va	65°	64°	· 74°
CA2	35°	24°	14°
X	v.p.gr.	cls.	v.p.gr.
Y	p.gr.	v.p.gr.	p.gr.
Ζ	bl.gr.	p.gr.	ol.gr.
D (g/cc)	3.082	3.058	3.084
(Å)	9.790	9.906	9.879
	17.959	18.007	18.091
20	5.291	5.279	5.290
3	104° 35'	104°24′	104° 38'
$a_{\alpha}\sin\beta$ (Å)	9.475	9.595	9.559
V (Å ³)	900.3	912.1	914.8

v. = very, p. = pale, gr. = green, bl. = blue, cls. = colorless and ol. = olive.

1. Richterite, Lake Saarinen. (R86-33.40)

2. Richterite, Lake Särkilampi. (87-1/KP-67)

3. Actinolite, east of Lake Särkilampi. (55-9/KP-67)

The X-ray data for the Siilinjärvi amphiboles are in agreement with the values given for synthetic minerals (cf. Kohn and Comeforo, 1955; Huebner and Papike, 1970).

Chemical composition

Table 2 gives the chemical composition of the Siilinjärvi amphiboles. The atomic ratios have been calculated on the basis of 23 oxygen atoms per unit cell, thus eliminating the possible amounts of F and Cl in the mineral (*cf.* Simonen and Vorma, 1969).

According to Leake (1968) richterite is defined as a Ca- and Al-poor, Na- and Si-rich amphibole of theoretical composition $Na_2CaMg_5Si_8O_{22}$ (OH)₂; consequently the name richterite may be used only when Si>7.25, Ca+Na+K>2.50, Ca < 1.50 and Na > 1.00. The values for the Siilinjärvi richterite are comparable with these values.

In addition to Na richterite also contains some K. A complete solid solution series between sodic richterite NaNaCaMg₅Si₈O₂₂(OH)₂ and potassic richterite KNaCaMg₅Si₈O₂₂(OH)₂ can be synthesized (Huebner and Papike, 1970). The relatively high content of K in the Siilinjärvi richterite is also in accordance with the presence of great amounts of phlogopite in the ultramafic intrusion phase of the complex.

In the Y-type cations of richterite a substitution of Mg by Fe²⁺ exists giving rise to ferrorichterite with up to 27.27 % FeO (Chapman and Williams, 1935). The replacement of Mg by Fe³⁺ is also common, the mineral in general having a high Fe³⁺: Fe²⁺ ratio (Deer *et al.*, 1965, p. 353); an extreme example is ferririchterite (Bilgrami, 1955) with 18.44% Fe₂O₃. Manganous richterite, up to 8.69 % MnO, has been found associated with manganese minerals (Sundius, 1946).

The Siilinjärvi actinolite from syenite contains more Na and K than the analyses given in the literature (cf. Deer et al., 1965, p. 251–253). In some tremolites moderate amounts of Na are

TABLE 2

Chemical analyses (Wt%) and atomic ratios based on 23 oxygen atoms per unit cell of richterite from glimmerite and actinolite from syenite, Siilinjärvi carbonatite complex. Analyses by P. Ojanperä

	1.	2.	3.
SiO	53 13	56.03	55.04
TiO	0.14	0.10	0.19
A1 O	2 13	0.10	1.50
Fe O	7 78	5.68	1.30
FeO	3 1 2	1 5 9	6.04
MnO	0.11	0.07	0.04
MgO	16.95	20.19	17 29
CaO	6.57	5.80	11 19
Na O	5.07	1.87	1 56
K O	1 58	2 62	0.68
P.O.	0.10	0.06	0.05
CO	0.72	0.00	0.05
$H_{-}O^{+}$	1.92	1 7 1	1 67
H_2O^-	0.04	0.05	0.03
1120	0.01	0.05	0.05
	99.66	99.63	99.60
이 일이다. 영화 영화		1	1
- (Si	7.65	7.85	7.79
2 [Al[⁴]	0.35	0.13	0.21
(Al[⁶]	0.01		0.04
Ti	0.02	0.01	0.02
\mathbf{v} Fe ³⁺	0.84	0.60	0.43
$1 Fe^{2+}$	0.37	0.19	0.72
Mn	0.01	0.01	0.04
Mg	3.62	4.21	3.64
[Ca	1.01	0.88	1.70
X, A {Na	1.41	1.32	0.43
[K	0.29	0.47	0.12
ΣΖ	8.00	7.98	8.00
$\Sigma Y \dots$	4.87	5.02	4.89
Σ X, A	2.71	2.67	2.25
OH	1.84	1.60	1.58
100 Mg	74.7	84.2	75.5
Mg + Fe + Mn			

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present and it is possible that there is a chemical series between tremolite and richterite (Sundius, 1946). The richterite formulae may namely be regarded as derived from tremolite by the substitution of Na for Ca in X accompanied by the addition of Na in the A-site.

Optical and physical properties

The optical and physical properties of the Siilinjärvi amphiboles are given in Table 1.

Pleochroism is usually very weak, bluish tints are, however, sometimes observed. The minerals, especially richterite, exhibit abnormal interference colors due to the strong dispersion of the bisectries. The weak pleochroism colors and the relatively low refractive indices of richterite may be essential for the optical identification of the mineral from other alkali amphiboles.

Discussion

Richterite occurs in various geologic environments such as in metamorphosed limestones (Sundius, 1946), potassic lavas (Carmichael, 1967), kimberlites (Erlank and Finger, 1970) and in carbonatites.

A wide variety of amphiboles has been reported in carbonatite complexes, ranging from ordinary hornblende and tremolite through somewhat sodic amphiboles to strongly alkaline types such as riebeckite (Heinrich, 1966). In East Africa richterite appears to be confined to calcite-rich rocks such as carbonatites, and has not been observed in fenites (Sutherland, 1969). In Lesnaya Varaka, Kola Peninsula, it is found in dolomite-amphibole rocks (Kukharenko *et al.*, 1965). The richterite from the carbonatites at Iron Hill, Colorado, has been described by Larsen (1942), who, like Miyashiro (1957), has used the name soda tremolite instead of richterite.

The occurrence of richterite in the glimmerite of the Siilinjärvi complex is in agreement with the general chemical character of the rocks. The presence of sufficient quantites of Na and Fe would have given rise to alkali amphiboles such as riebeckite and arfvedsonite, as is generally the case in carbonatite complexes. This also matches the statement of Ernst (1968) that richterite is evidently restricted to high temperature environments and to rock bulk compositions rich in both Na and Ca relative to Al.

Author's note

After this manuscript was submitted to press, an error was discovered in the K—Ar age determinations of the Siilinjärvi minerals (cf. Puustinen, 1971, p. 38). The computer calculations showed too large argon contents. The corrected values are as follows:

Mineral	Age (m. y.)	Ref. No.
Phlogopite	 2030 ± 30	204—2/KP—69
»	 1785 ± 30	85—1/KP—67
>>	 1790 ± 30	46—8/KP—67
>>	 1985 ± 30	71—6/KP—67
Richterite	 2530 ± 45	71—5/KP—67
Actinolite	 2260 ± 42	55—9/KP—67

Acknowledgements — Dr. Atso Vorma, Geological Survey of Finland, gave the author valuable aid in the X-ray determinations and he also critically read the manuscript. Chemical analyses were made by Mr. P. Ojanperä, M. A., Geological Survey of Finland. Mr. John Nelson, M. Met., Technical University, Otaniemi, checked the language of the English manuscript.

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Manuscript received, June 10, 1971.